Multi-scale Hydrological **Data Assimilation in** Layered Media Juan M. Restrepo **Department of Mathematics Physics Department** University of Arizona

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## TIME SERIES Estimation Problems:

Given a random time series { z(t): t < t<sub>0</sub>} z(t) 2  $\mathbb{R}^{N}$ 

Prediction:
 Estimate {z(t): t> t<sub>0</sub>}
 Filtering (Nudiction):
 Estimate {z(t<sub>0</sub>)}
 Smoothing (Retrodiction):
 Estimate {z(t): t · t<sub>0</sub>}

#### Turning a model into a state estimation problem Example: $\partial_t u(z,t) = v \partial_{zz} u(z,t) + f(t)$ $u(z,0) = u_0(z)$ $u(0,t) = g(t) \quad u(1,t) = h(t)$ Discretizing: $x(t) \in [u_1(t), u_2(t)...u_N(t)]^T$ is the state variable, obeying $x(t+\delta t) = A x(t) + B q(t)$ $x(t) = A x(t-\delta t) + B q(t-\delta t)$ Leads to LINEAR PROBLEM: $L(x(0),...,x(t-\delta t),x(t),x(t+\delta t),...,x(t_{f}),...,x(t_{f}))$ $Bq(t), Bq(t+\delta t), \dots, t) = 0$

 $x(t) 2 \mathbb{R}^N$  Bq2  $\mathbb{R}^N$ 

#### **Statement of the Problem**

MODEL (Langevin Problem):  $dx(t) = f(x(t), t)dt + (2D)^{1/2}(x, t)W(t), \qquad t > t_0,$   $x(t_0) = x_0.$   $x, f, dW \in \mathbb{R}^N,$ 

#### DATA:

$$egin{aligned} y(t_m) &= h(x_m) + [2R(x_m,t)]^{1/2} \epsilon_m \ & ext{where } m = 1,2,...,M \ & ext{h}, \epsilon &: & oldsymbol{R}^N o oldsymbol{R}^{N_y} \end{aligned}$$

#### GOAL: estimate moments

(at least) find mean conditioned on data:  $x_{s}(t) = E[x(t)|y_{1},...,y_{M}]$ and Covariance matrix (uncertainty)  $C_{s}(t) = E[(x(t)-x_{s}(t))(x(t)-x_{s}(t))^{>}|y_{1},...,y_{M}]$ 

The conditional mean  $x_s(t)$  minimizes tr  $C_s(t) = E[|(x(t)-x_s(t))|^2|y_1,...,y_M]$ . It is termed the smoother estimate.

#### Optimal Estimate of Discretized Linear Model with Gaussian Noise

Let  $z_i = u(x_i)$  where  $x_i 2 \Omega$   $B z + n_m = F$   $D z + n_d = Y$  OR M z + N = T $min_z J = \langle N^T N \rangle$ 

Least Squares, SVD (Kalman)

#### **A Nonlinear Example**

Stochastic Dynamics (Langevin Problem):  $dx(t) = f(x(t)) dt + \kappa dW(t)$ with  $V(x) = -2x^2 + x^4$   $f(x) = -V'(x) = 4x(1-x^2)$  $\kappa = 0.5$ 

Measurements:

at times m  $\Delta t$ , m=1,..., M one observes  $y_m := X(t_m) + \rho N_m$ to have measured values  $Y_m$ , m=1,2,...,M Kolmogorov Equation

$$\partial_t P = -\partial_x [f(x) P] + \kappa^2 \partial_{xx} P/2$$

$$P(x,t)_{t!1} = P_s(x)$$



#### Observations



#### **BAYESIAN STATEMENT**

- P(X|D) / Prior £ Likelihood
- Use data for the likelihood
- Use model for the prior

 $P(X|D) \sim exp(-A_{data}) exp(-A_{model})$ 

#### **Extended Kalman Filter**



## **Alternative Approaches**

- KSP: optimal, but impractical
- ADJOINT/4D-VAR: optimal on linear/Gaussian

(Restrepo, Leaf, Griewank, SIAM J. Sci Comp 1995)

Mean Field Variational Method

(Eyink, Restrepo, Alexander, Physica D, 2003)

- enKF (ensemble Kalman Filter)
- Particle Method

(Kim Eyink Restrepo Alexander Johnson, Mon. Wea. Rev. 2002)

#### Path Integral Method

(Alexander Eyink Restrepo, J. Stat. Phys. 2005 and Restrepo Physica D, 2007)

## Path Integral Method

- Related to simulated annealing
- It could be developed as a black box
- Simple to implement
- Can handle nonlinear/non-Gaussian problems
- Calculates sample moments

PROBLEM: Relies on MC!!!

$$dx(t) = f(x(t), t)dt + [2D(x, t)]^{1/2}dW(t), \qquad t > t_0,$$
  
$$x(t_0) = x_0.$$

Discretized using explicit Euler-Maruyama scheme

$$\begin{array}{ll} x_{k+1} &=& x_k + f(x_k, t_k) \delta t + (2D)^{1/2} (x_k, t_k) (W(t_k + \delta t) - W(t_k)), \\ & k = 0, 1, 2, \ldots \end{array}$$

 $x_{k=0} = x_0.$ 

Let 
$$\eta(t_k) = W(t_k + \delta t) - W(t_k)$$
,  
at times  $t_k$ ,  $k=0,1,2,...,$ 

Suppose  $\eta(t_k)$  is Gaussian Prob  $\eta(t) \approx \exp(-1/2 \sum_k | \eta(t_k) |^2)$ .

Hence exp(- $A_{dyn}$ ), for t = t<sub>0</sub>, t<sub>1</sub>, ...t<sub>T</sub>

$$\begin{array}{l} A_{dyn} \quad \sum_{k = 0}^{T-1} \left[ \left[ (x_{k+1} - x_k) / \delta t - f(x_k, t_k) \right]^{>} D^{-1}(x_k, t_k) \\ \left[ (x_{k+1} - x_k) / \delta t - f(x_k, t_k) \right] \right] / 4 \end{array}$$

To include influence of observations use Bayes' rule. This modifies Action:

 $A_{obs} = \sum_{m=1}^{M} [h(x(t_m) - y(t_m))] R^{-1}[h(x(t_m)) - y(t_m)]$ 

The Total Action:

$$A = A_{dyn} + A_{obs}$$

The Action is like the log-likelihood.

#### **PIMC Filter Results**





# Estimation Applied To Steady State Hydrology



- Estimate hydraulic head in domain
- Estimate material properties in domain
- Estimate "best"
   boundary values
- Estimate all of the above

Simplest Boundary Value Problem

 $n(x),\, heta(x),\, 
ho(x)\,$  are known statistical quantities

#### OUR APPROACH

 USE DATA-DRIVEN CLASSIFICATION: estimates partitioning into homogeneous layers.
 Support Vector Machines

DISCRETIZE Variational formulation for the model plus constraint (via Lagrange multiplier): constrained minimum satisfies E-L. Coupling of each subproblem is automatically satisfied.

Weak form (using Dirichlet energy)

SOLVE nonlinear system in each subdomain: Newton

#### **Data-Driven Classification**

Estimate the boundaries between heterogeneous geologic facies

- Data
  - $K_i = K(\mathbf{x}_i)$ , e.g., conductivity  $h_{jk} = h(\mathbf{x}_j, t_k)$ , e.g., head
- Data are sparse
- Measurements are well differentiated



Measurements of system parameters  $(K) \implies$  forward FD problem Measurements of system states  $(h) \implies$  inverse FD problem Assign indicators to data,

 $I(\mathbf{x}_i) = 1(0)$  if  $\mathbf{x}_i \in M_1(M_2)$ 

- $\mathcal{I}(\mathbf{x}, \boldsymbol{\alpha}) \equiv$  an estimate of  $I(\mathbf{x})$
- min  $R = \int ||I \mathcal{I}|| \mathrm{d}P(I, \{\mathbf{x}\}_{i=1}^N)$
- Geostatistics (Kriging)
  - 1. the  $L^2$  norm
  - 2. the indicator function  $I(\mathbf{x})$  is a random field, and
  - 3. the choice of sampling locations  $\{\mathbf{x}_i\}_{i=1}^N$  as deterministic.  $\implies$

4. Variance: 
$$\sigma_I^2 = \int (I - \mathcal{I})^2 dP(I)$$

#### SVMs

- 1. the  $L^1$  norm
- 2. the indicator function  $I(\mathbf{x})$  as deterministic, and
- 3. the choice of sampling locations  $\{\mathbf{x}_i\}_{i=1}^N$  as random.  $\Longrightarrow$
- 4. Expected risk: min  $R_{exp} = \int |I \mathcal{I}| dP(\{\mathbf{x}\}_{i=1}^N)$



## **Support Vector Machines**

- Alternative to Kriging
- Very good alternative when sample densities are too low for Kriging
- Highly automated
- Can be incorportated in the solver problem

#### Heterogeneous Sub-Surface

In each subdomain i = 1, 2, ..., M

$$K(x,\omega) = \exp\left[\sum_{j=1}^{\infty} \kappa_j(\omega)\phi_j(x)\right]$$
$$u(x,\omega) = \sum_{j=1}^{\infty} \mu_j(\omega)\phi_j(x)$$

$$egin{aligned} -
abla \cdot (K
abla u) &-\overline{f} = n(x,\omega) \ E(n) &= 0 \ E(n(x)n(y)) &= g(|x-y|) \end{aligned}$$

## (Weak) Variational Formulation

Let P:=[u,K]

- Use standard machinery to solve nonlinear problem but use weighted norms (locally in each subdomain).
- Use Newton solver but decide whether to do global estimate of partial estimates (increasing or decreasing the uncertainty in each subdomain).

Use Galerkin discretization of Newton Systems.

## Weak Formulation (no noise) $\phi(P) = \frac{1}{2} ||T(P) - y||^2 + S(P - P_0)$

Dirichlet-like Energy  $S(P) = \sum_{i=1}^{M} \left[ \int_{\Omega_i} \left( \frac{1}{2} |\nabla P|^2 + kP^2 \right) dx \right]$ 

$$G(P) = -\nabla \cdot (K\nabla P) - \overline{f} = 0$$
  

$$\Phi(P, \Lambda) = \phi(P) - \langle \Lambda, G(P) \rangle$$
  
LEADS TO: Find [ P,  $\Lambda$  ]<sup>T</sup> such that  

$$\langle \Phi'(P), v \rangle - \langle \Lambda, G'(P)v \rangle = 0, \quad \forall v \in X$$
  

$$\langle \nu, G(P) \rangle = 0, \quad \forall v \in Y^*$$

X, Y\* Banach spaces

## Newton Solution Find corrections $[\pi, \lambda]^{\mathsf{T}}$ to $[\mathsf{P}, \Lambda]^{\mathsf{T}}$ $\begin{bmatrix} \phi''(P) - [G''(P)^T \Lambda] & -G'(P)^T \\ G'(P) & 0 \end{bmatrix} \begin{bmatrix} \pi \\ \lambda \end{bmatrix} = \begin{bmatrix} -\phi'(P) + G'(P)^T \Lambda \\ -G(P) \end{bmatrix}$

To find Hessians and Jacobians, use ADIFOR/C

#### **Final Comments**

- Model error formulation vs. closure?
- Already existing nonlinear solvers.
- Weak formulation automatically takes care of boundary conditions at the layer interfaces.
- Can give a-priori estimates of error.
- Unlike Inverse Method (Tikhonov, e.g.) problem is greatly more numerically stable.
- Use PIMC (see Restrepo, 2007) to benchmark results.
- Constrain number of SVM subdomains to the Newton solve.

#### **Further Information:**

http://www.physics.arizona.edu/~restrepo